GM 59161

A REPORT ON THE GRENVILLE VIII PROPERTY

A Report on the Grenville VIII Property

For: **Christian Derosier, prospector**

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Introduction

During August and September, 2000 three (3) related exploration programs were completed in the Grenville-Kilmar area by Christian Derosier and the writer under the Québec Mining Exploration Assistance Program. The first was a "Volet A1" regional prospecting program. The other two were "Volet A2" property-specific exploration programs, including the one described in this report – the Grenville VIII Project.

The objective of the Grenville VIII exploration program was to assess the potential of the Grenville VIII Property and immediate surroundings to host economic concentrations of magnesite, wollastonite, and calcite within the crystalline limestone sequence that apparently underlies most of the property. The exploration methodology employed included prospecting and sampling bedrock exposures, completion of magnetometer, geological mapping, and major element analyses of selected samples.

Despite the existence of the former Kilmar magnesite mine 4 km NNW of the Grenville VIII Property along the same broad sequence of metasediments and crystalline limestones, no previous exploration is recorded on the Grenville VIII Property. A hydrothermal origin to the magnesite at Kilmar is proposed and is pertinent to exploration in this sequence.

Property Location and Description

The Grenville VIII Property consists of three (3) contiguous claims located in Range VIII of Grenville Township, Argenteuil County, Québec. The three claims correspond to lots 9 (claim 5233595), 10 (claim 5233596), and 11 (claim 5233597) each covering a surface area of 110 hectares for a total of 330 hectares. The claims are presently in good standing until February 20, 2002.

The Grenville VIII Property is located ten (10) km north of the town of Grenville, Québec, which is connected by bridge across the Ottawa River to the town of Hawkesbury, Ontario. The Canadian National Railway line, along the north side of the Ottawa River, passes nine (9) km south of the property. The center of the property corresponds to UTM coordinates 5063240 N, 531275 E.

The property is accessed from provincial Highway 148 via the gravel Scotch Road which continues northward out of Grenville and traverses the eastern and northern portions of the property ending at the former Kilmar Road. The paved Kilmar Road passes two (2) km west of the property. Several usable bush roads access most of the property. The bed of the railway which formerly connected the Kilmar Mine to the CNR line and is now used as an ATV trail, traverses the northeast part of the property.

A small round lake is situated in the north-central part of the property. It drains into the Calumet River, which is just a broad, shallow stream here, and flows southward across the southwest part of the property. This area is flat and covered by thick fluviatile sand and gravel deposits. The eastern portion of the property is comprised of low hills of glacial till and outcrop, interspersed with flat sandy and swampy areas. High rocky hills occur in the northwest corner and in the south-central part of the property.

Forest cover is mixed temperate with beech, maple, birch, poplar, spruce, hemlock, and white pine. Selective forest harvesting is carried out in this area. Elevations range from 230 to 250 m over a broad, flattish area in the central, north-central, southwest, and eastern parts of the property, to 370 m. in the northwest corner, and 340 m on the pronounced hill in the south-central area.

The Hydro-Québec power grid is accessible three (3) km to the northwest just south of the former Kilmar Mine.

Property History and Previous Work

No previous exploration is recorded on or in the immediate vicinity of the Grenville VIII Property. Considering the proximity of the Kilmar magnesite deposits, four (4) km to the NNW, and the lengthy, albeit interrupted, period of mining and exploration there (1900 to 1992), this lack of exploration around Grenville VIII is probably due more to deep overburden and failure to find significant magnesite showings and boulders than anything else.

Geology

Regional Geology

This region was initially mapped by Sir William Logan in the 1856. In 1902 Ells published a 1 inch = 4 miles map of the area (G.S.C. Grenville sheet, map 750). M. E. Wilson mapped the area in 1916 at a 1 inch = 1 mile scale (G.S.C. map 1680). Subsequently the region was mapped at 1 inch = 1mile scale by F. F. Osborne and H. W. McGerrigle in 1936 (Rapport Annuel du Service des Mines de Québec pour l'année 1936, partie C – Région de Lachute). This map still forms the best geological base for the region. Grenville Township was mapped in 1976 by A. R. Philpotts (M.R.N. rapport géologique 156) as part of an M. Sc. Thesis.

Logan coined the term Grenville Series in reference to the metasedimentary sequence composed of crystalline carbonates, quartzites, and garnetiferous quartzo-feldspathic gneisses occurring in broad, roughly north-south trending swathes in this area. The Grenville Series metasediments are bounded by more extensive igneous masses. The metasediments are usually highly contorted. However, gently folded crystalline carbonates occur over significant areas. The mass effect is important in this regard. A thick, fairly uniform mass of carbonate rock is less susceptible to folding than a mixed sedimentary sequence of corresponding thickness.

The crystalline carbonates weather recessively in comparison to the quartzites, quartzo-feldspathic gneisses, and bounding igneous rocks. Extensive areas of thick glacial overburden occur throughout the region. Examples include the Lake McDonald and Harrington-Rivington areas to the north, and the Avoca area to the west. These areas are underlain predominantly by carbonate-dominant metasediments. Less extensive examples include the area adjacent and to the south of the Kilmar Mine, large parts of the Grenville VIII Property, and the Grenville IV Property area.

Following deposition and folding of the metasediments, the bounding, regionally predominant igneous intrusions were emplaced. Compositions of them range from gabbroic through monzonitic, mangeritic, granitic and syenitic. They occur as large, elongate masses roughly paralleling the regional north-south trend, and as rounded, batholithic masses. Gneissosity is developed locally within them. The igneous rocks are collectively referred to as the Morin Series.

A set of east-west trending diabase dykes cut all of the above units. They are up to 50 m thick.

A thin layer of Pleistocene age boulder till covers most of the elevated portions of the region. The valleys are filled with highly variable thicknesses of fluvio-lacustrine deposits of sand, gravel, and clay.

The crystalline limestones commonly contain abundant inclusions of paragneiss, pegmatite, and feldspar porphyroblasts. They have rough weathered surfaces. On fresh surface they are medium-grained, whitish-grey rocks composed largely of calcite, with minor, variable amounts of graphite, green diopside, and serpentine. The crystalline limestones in this region usually host disseminated graphite grains. The graphite occurs as evenly distributed, discrete euhedral grains and commonly forms 2 to 3%, but may form up to 8 to10% of the rock. Green diopside is another fairly ubiquitous, but minor, component of them.

Locally, abundant inclusions of rounded blocks of paragneiss within the crystalline limestones give the rock a conglomeratic aspect. The possible origins for them are problematic. They may be the rolled up, metamorphosed equivalents of small layers of sediment within the limestones. Or they may be metamorphosed olistostromes.

Thermal metamorphism of the crystalline limestones occurs proximal to Morin Series igneous rocks. Sporadic minor occurrences of wollastonite with abundant green diopside are evidence of this (GR-139). Thermal metamorphism has locally caused the limestones to be recrystallized to coarse-grained dolomites.

Locally, the crystalline limestones contain abundant phlogopite, muscovite, and/or tremolite. Good examples of this occur in roadcuts along the Kilmar Road up to five (5) km south of the former Kilmar Mine (eg. sample GR-120), and at the north end of the Grenville VIII Property (GR-67). The micas are alteration products indicative of potassic hydrothermal activity. Closer to the mine the limestones are completely altered by serpentinization and introduction of sulphides. South of the Kilmar Mine sulphide hydrothermalism has resulted in the deposition of elevated concentrations of copper, lead, zinc, and silver (GR-121). Another exposure exhibiting serpentinization and sulphide hydrothermalism occurs along the Harrington Road just west of Highway 327 (GR-136), three (3) km WNW of the former mine. **Thus a metallogenetically significant episode of hydrothermal activity becomes evident in this area, which includes the Grenville VIII Property.**

Property Geology

The Grenville VIII Property lies on a broad, locally interrupted sequence of Grenville Series metasediments, which extends over 20 km to the north and south to the Ottawa Graben. The former Kilmar magnesite mine is situated four (4) km NNW of the property.

Almost half of the property is flat-lying and is covered by fluviatile sand and gravel deposits, ice-contact stratified drift, and swampy peat bogs. This type of terrain includes all of the central, north-central, and southwest areas, as well as portions of the east and southeast. The low hills in the northeast, east (Lac Noir area) and southeast (Lac Cook area) are predominantly formed of crystalline limestone.

Northwest of Lac Noir, along and west of the old rail bed, a large area of massive, buff-coloured, gently-folded crystalline limestones occurs (GR 53, 54, 59, 60, 62, 63). They contain 2-4% disseminated graphite but, otherwise, appear to be relatively pure. A thin layer of glacial till covers 99% of this area. The elevated area to the northeast and east of these limestones is biotite-feldspar-quartz paragneiss (GR 49, 50, 51, 61)

The area west of Lac Noir extending south to Lac Cook is predominantly crystalline limestone (GR 23, 24, 25, 26). However the magnetic response over these areas is inconsistent with a massive, uniform limestone and folding with subordinate quantities of interbedded, subcropping paragneisses is suspected.

The prominent high hill in the south-central part of the property is formed of quartzites along the north and west flanks (GR 64) with a WNW gneissosity. The west slope of it is predominantly a mylonitic quartz-feldspar gneiss (GR 65, 66). Locally, this rock has the aspect of a sheared rhyolitic horizon. The peak of this hill is a granitic gneiss, possibly with a granite origin. Outcrops 500 to 750 m south of the south boundary support a volcanic history for the south-central and southwest parts of the Grenville VIII Property. GR 21 is a silicified volcanic rock with a tuffaceous aspect. GR 22 is a banded gneiss resembling a rhyolite. GR 48 is a brecciated felsic to intermediate volcanic tuff with a carbonate-rich matrix and over 5% sulphide content. Due to weathering and the friability of this occurrence it was not possible to take a fresh sample for assay.

The high hills to the west of the Calumet River and in the northwest portion of the property are predominantly quartz-feldspar-biotite paragneisses. The writer was unable to examine these adequately and they deserve further attention. Three suspected outcrops were found in the sandy overburden of the Calumet valley (GR 55, 56, 57). All were flat, small, and difficult to sample. GR 55 appears to be either a breccia or a volcanic fragmental and is formed of siliceous, angular fragments in a crystalline matrix. GR 56 exhibits siliceous layering with a westerly strike.

At the north limit of the property, where the Scotch Road curves toward the west, an important outcrop (GR 67) of magnesite-tremolite-quartz-(muscovite?) rock occurs. This assemblage is suggestive of a hydrothermally altered dolomite. Unfortunately due to the almost total coverage of this part of the property by stratified drift and swampy areas no additional outcrops were found in this area.

The magnetic survey covers approximately 60% of the property. Hazards imposed by illegal marijuana cultivation resulted in inadequate coverage in the north and south-central areas. However, the survey indicates that a north-south contact zone, which roughly follows the baseline, separates a carbonate domain to the east from quartzitic/felsic gneiss terrain to the west. Thusfar geology supports this division. The contact zone is a series of elongate, intense magnetic highs and lows parallel to, and just west of, the baseline.

The magnetic susceptibility of the carbonates is not uniform due to minor, folded more mafic gneiss horizons within the carbonates and/or variable overburden thicknesses. Metasomatism of the carbonates may also be a factor.

West of the contact zone, survey coverage is inadequate. However, the few samples from this area indicate a quartzitic and felsic gneiss dominant terrain. In the south-central-southwest area, these lithologies are interpreted by the writer to have a volcanic derivation. Inadequate geological and magnetic coverage in the northwest preclude any interpretation of that part of the property until further work can be done.

Conclusions and Recommendations

Geological interpretation of the south-central-southwest area of the property suggests a felsic volcanic origin for the siliceous and felsic gneisses here. This terrain extends to the south of the west part of the property for a kilometre or more. Further geological and magnetic investigation is needed to extend this terrain toward the north, onto the northwest portion of the property.

Approximately the eastern half of the property is underlain by crystalline limestone and dolomite. Most of this appears to be gently folded and massive. However, magnetic response suggests some combination of tighter folding, minor mafic interbedding, and variable overburden thicknesses. Alteration due to metasomatic activity also affects the magnetic susceptibility, particularly approaching the contact zone.

The contact zone is a series of intense, elongate magnetic highs and lows situated just west of the baseline. It appears to extend the north-south length of the property, but magnetic coverage must be extended in the south-central area for better definition. This zone is interpreted to represent the contact between the volcanically-derived terrain to the west and the carbonate terrain to the east. No outcrops occur within the contact zone. The magnesite/tremolite content of sample GR 67, which is about 300 m east of the contact zone, indicates intense Mg hydrothermalism. Over 5% sulphides occur in a faulted felsic fragmental 750 m south of the south boundary (GR 48) on the flank of the interpreted southern extension of the contact zone.

It is postulated that the contact zone represents concentrated hydrothermal activity characterized by Mg metasomatism in the overlying carbonates and introduction of sulphides as local accumulations along the zone.

Consequently, the following exploration program is strongly recommended:

1) Complete the ground magnetic survey coverage in the northwest and southern parts of the property. Fill in the current line spacing to 100 m intervals.
2) Horizontal loop EM and VLF surveys should be completed the length of the property from L5+00E to L5+00W at the recommended 100 m line spacing.
3) Complete a thorough geological mapping and sampling program.
4) Drill the best targets interpreted from integration of results from steps 1, 2, and 3. The targets are sulphide accumulations along the contact zone and massive magnesite lenses in the carbonates.

Consideration should be given to staking the three lots adjacent to the south of the Grenville VIII Property (the lots to the north having already been staked).

J. D. Charlton, APGGQ, FGAC, ~~APGO~~
December 14, 2000

GÉOLOGUE GEOLOGIST
APGGQ
JOHN D. CHARLTON
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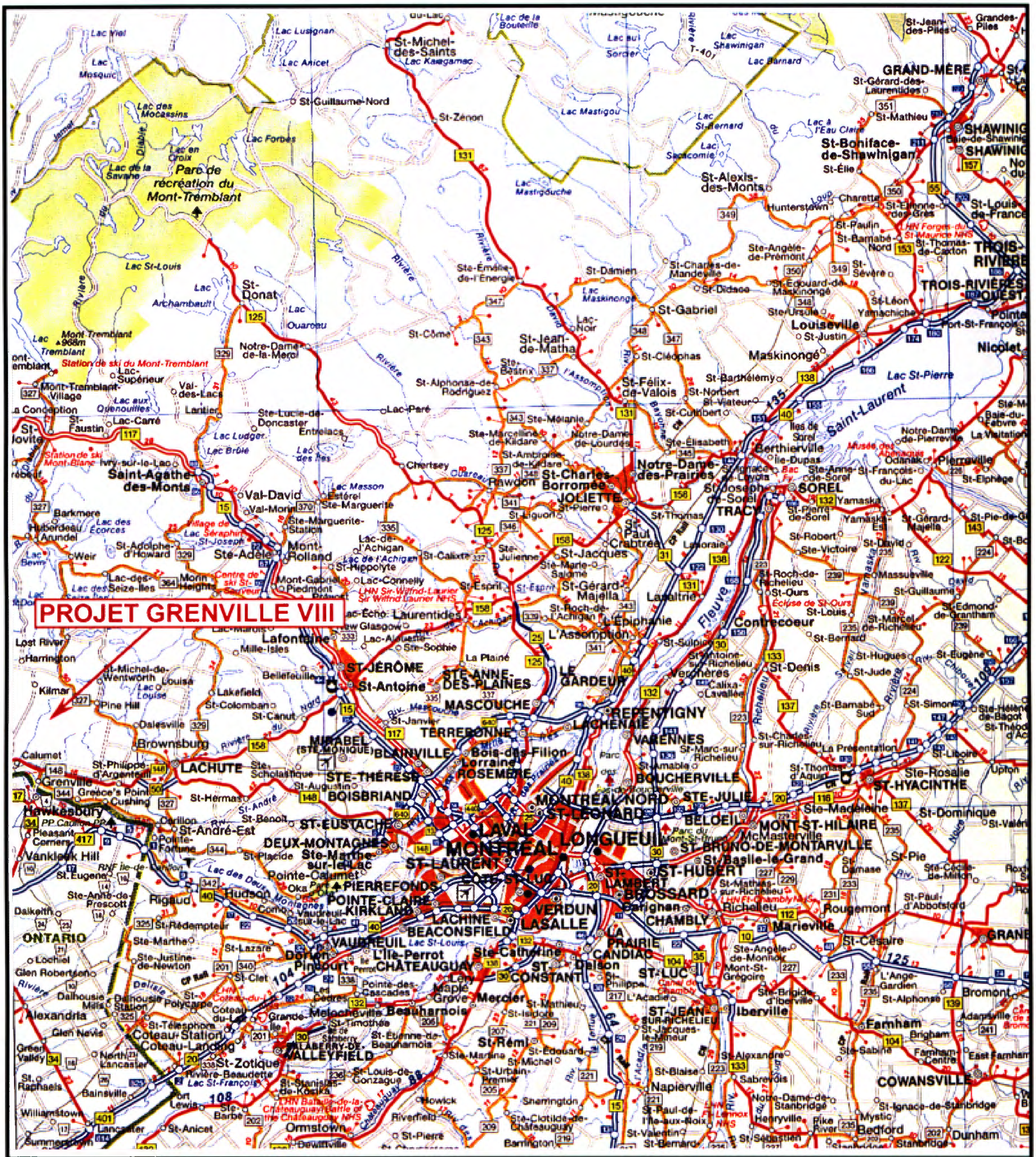
References

1) **Geosig Inc**., 2000, Levés RDTGPS, de magnétométrie et TBF (VLF) Propriété Grenville IV; Implantation d'une grille et levé magnétique Propriété Grenville VIII, Canton de Grenville, Québec.

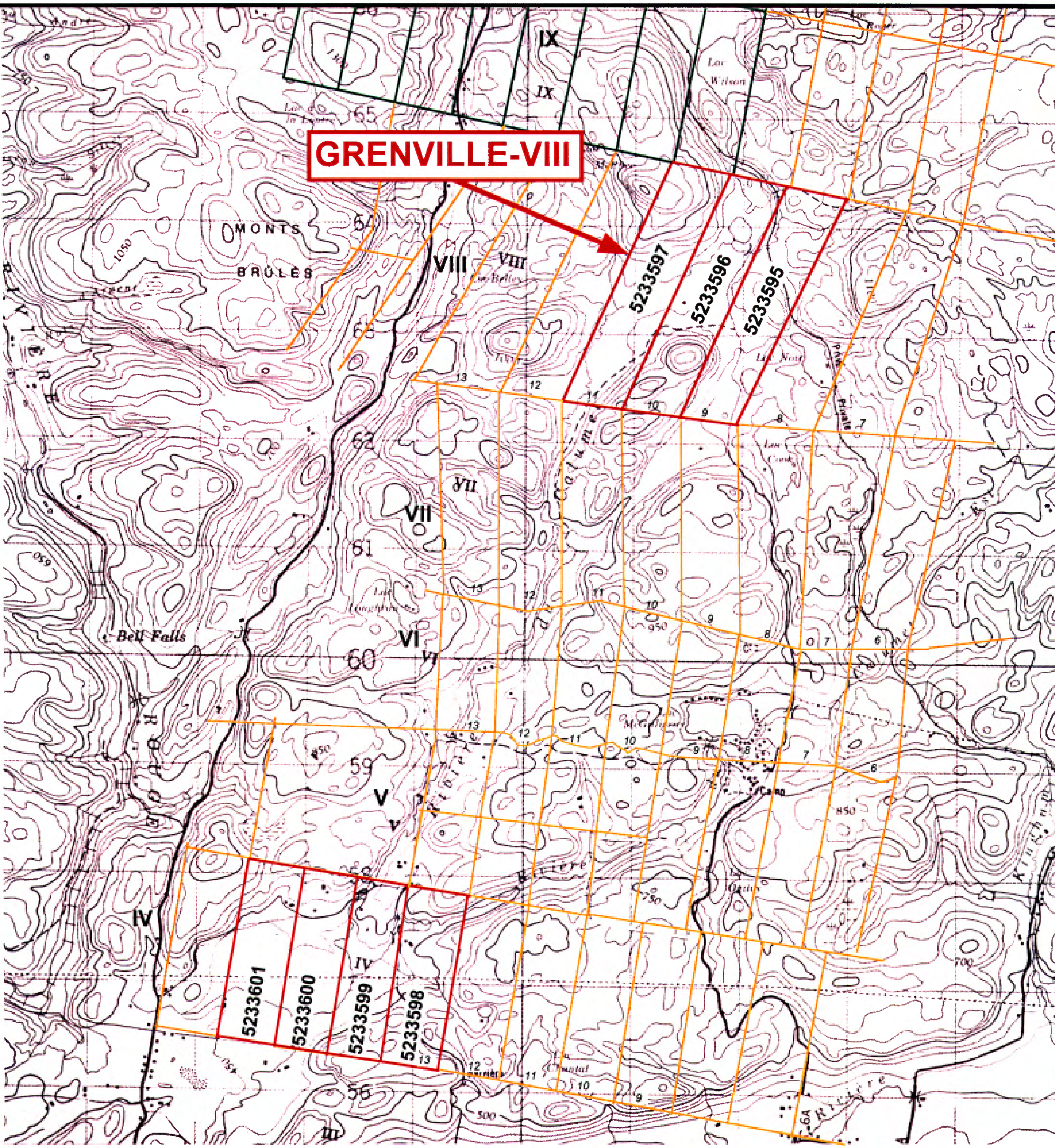
2) **Osborne, F. F. and McGerrigle, H. W**., 1936, Rapport annuel du Service des Mines, 1936, Partie C, Région de Lachute.

3) **Philpotts, A. R.,** 1976, Grenville County, SE part, Geological Report 156, Ministère des Richesses Naturelles.

4) **Richard, S. H.,** 1984, Surficial Geology, Lachute-Arundel, Quebec-Ontario; Geological Survey of Canada, Map 1577A, scale: 1:100,000.



CLIENT	CHRISTIAN DEROSIER & ASSOCIÉS	FAIT/ MADE J.D. CHALTON	DATE Décembre 2000
PROJET/ PROJECT	PROJET GRENVILLE VIII	APPR.	SCALE/ ECHELLE 1: 800 000
	CARTE DE LOCALISATION	CONTR.	SUBDIV. / No 1



ÉCHELLE 1: 50 000 PARTIE DE LA CARTE 31G/ 10

CLIENT CHRISTIAN DEROSIER & ASSOCIÉS PROJET/ PROJECT **PROJET GRENVILLE VIII**	FAIT/ MADE **J.D. CHALTON**	DATE **Décembre 2000**	
	APPR.	SCALE/ ECHELLE 1: 50 000	
CARTE DE LOCALISATION DES CLAIMS	CONTR.	SUBDIV.	No **2**

RAPPORT: C00-64122.0 (COMPLET) RÉFÉRENCE: 175037

CLIENT: CHRISTIAN DEROSIER SOUMIS PAR: C.DEROSIER
PROJET: AUCUN DATE RECU: 22-NOV-00 DATE DE L'IMPRESSION: 6-DEC-00

DATE APPROUVÉ	COMMANDE	ÉLÉMENT	NOMBRE D'ANALYSES	LIMITE INFÉRIEURE DE DETECTION	EXTRACTION	MÉTHODE
001210	1 SiO2	Silica (SiO2)	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	2 TiO2	TiO2 - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	3 Al2O3	Al2O3 - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	4 Fe2O3	Fe2O3*- IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	5 MnO	MnO - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	6 MgO	MgO - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	7 CaO	CaO - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	8 Na2O	Na2O - IC80	7	0.01 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	9 K2O	K2O - IC80	7	0.05 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	10 P2O5	P2O5 - IC80	7	0.03 PCT	FUSION BORATE	INDUC. COUP. PLASMA
001210	11 LOI	Perte au feu - IC80	7	0.05 PCT	Perte au feu 1000 C	GRAVIMETRIE
001210	12 Total	Wh Rock Total - IC80	7	0.01 PCT		
001210	13 Ba	Ba - IC80	7	10 PPM	FUSION BORATE	INDUC. COUP. PLASMA
001210	14 Cr	Cr - IC80	7	10 PPM	FUSION BORATE	INDUC. COUP. PLASMA
001210	15 Sr	Sr - IC80	7	5 PPM	FUSION BORATE	INDUC. COUP. PLASMA
001210	16 Ag	Ag - IC01	7	0.2 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	17 Cu	Cu - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	18 Pb	Pb - IC01	7	2 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	19 Zn	Zn - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	20 Mo	Mo - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	21 Ni	Ni - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	22 Co	Co - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	23 Cd	Cd - IC01	7	0.2 PPM	HCL:HNO3 (3:1).	INDUC. COUP. PLASMA
001210	24 Bi	Bi - IC01	7	5 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	25 As	As - IC01	7	5 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	26 Sb	Sb - IC01	7	5 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	27 Fe	Fe - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	28 Mn	Mn - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	29 TE	Te - IC01	7	10 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	30 Ba	Ba - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	31 Cr	Cr - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	32 V	V - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	33 Sn	Sn - IC01	7	20 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	34 W	W - IC01	7	20 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	35 La	La - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	36 Al	Al - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASMA
001210	37 Mg	Mg - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	38 Ca	Ca - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	39 Na	Na - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	40 K	K - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	41 Sr	Sr - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	42 Y	Y - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	43 Ga	Ga - IC01	7	2 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	44 Li	Li - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	45 Nb	Nb - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	46 Sc	Sc - IC01	7	5 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	47 Ta	Ta - IC01	7	10 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	48 Ti	Ti - IC01	7	0.010 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	49 Zr	Zr - IC01	7	1 PPM	HCL:HNO3 (3:1)	INDUC. COUP. PLASM
001210	50 S	S - IC01	7	0.01 PCT	HCL:HNO3 (3:1)	INDUC. COUP. PLASM

TYPES D'ÉCHANTILLONS	NOMBRE	FRACTION UTILISÉE	NOMBRE	PRÉP. DE L'ÉCHAN.	NOMBRE
ROCHE	7	-150	7	CONCASSER, PULVERISE	7

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CHIMITEC
BONDAR CLEGG

Rapport Lab Geochimie
Geochemical Lab Report

CLIENT : CHRISTIAN DEROSIER

PROJET: AUCUN

RAPPORT: C00-64122.0 (COMPLET)

DATE RECU : 22-NOV-00 DATE DE L'IMPRESSION: 6-DEC-00 PAGE 1A(1/ 2)

NUMÉRO DE L'ÉCHANTILLON	ÉLÉMENT UNITÉS	SiO2 PCT	TiO2 PCT	Al2O3 PCT	Fe2O3 PCT	MnO PCT	MgO PCT	CaO PCT	Na2O PCT	K2O PCT	P2O5 PCT	LOI PCT	Total PCT	Ba PPM	Cr PPM	Sr PPM
26307		54.72	0.07	1.01	2.06	0.07	16.59	24.02	0.26	0.20	<.03	0.17	99.21	33	64	214
26308		22.17	0.25	4.48	2.09	0.02	1.53	38.69	0.15	1.28	0.04	27.93	98.82	239	142	1527
26312		59.79	0.10	0.76	1.67	0.07	14.84	20.33	0.30	0.20	0.07	0.51	98.70	298	85	283
26313		8.45	0.11	1.54	0.62	0.01	0.84	46.79	0.25	0.61	<.03	37.77	97.44	561	39	3914
26315		10.46	0.11	2.20	0.76	<.01	0.96	45.51	0.42	0.86	0.05	35.87	97.68	698	33	4102
26316		78.29	0.17	2.30	0.85	0.03	0.65	13.04	0.71	0.76	0.06	2.03	98.93	74	231	181
26317		7.23	0.10	1.15	0.60	<.01	0.89	48.11	0.07	0.46	0.05	38.36	97.51	461	51	4401

NUMÉRO DE L'ÉCHANTILLON	Ag PPM	Cu PPM	Pb PPM	Zn PPM	Mo PPM	Ni PPM	Co PPM	Cd PPM	Bi PPM	As PPM	Sb PPM	Fe PCT	Mn PPM	TE PPM	Ba PPM	Cr PPM	V PPM	Sn PPM	W PPM	La PPM	Al PCT	Mg PCT
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26308	0.3	10	<2	19	<1	12	5	<.2	<5	<5	<5	1.36	102	<10	6	64	11	<20	<20	12	.42	.44
26312	<.2	4	<2	12	2	4	2	<.2	<5	<5	<5	0.24	70	<10	110	43	3	<20	<20	3	.07	.89
26313	0.7	4	<2	5	<1	4	2	<.2	<5	<5	<5	0.42	93	<10	15	14	4	<20	<20	6	.14	.22
26315	0.8	5	2	5	<1	6	2	<.2	<5	<5	<5	0.56	56	<10	16	19	3	<20	<20	7	.10	.29
26316	<.2	4	<2	2	<1	5	2	<.2	<5	<5	<5	0.30	142	<10	4	127	4	<20	<20	15	.11	.06
26317	0.7	4	<2	3	<1	4	1	<.2	<5	<5	<5	0.33	43	<10	47	11	1	<20	<20	3	.21	.09

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Rapport Lab Geochimie
Geochemical Lab Report

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NUMÉRO DE L'ÉCHANTILLON	ÉLÉMENT UNITÉS	Ca PCT	Na PCT	K PCT	Sr PPM	Y PPM	Ga PPM	Li PPM	Nb PPM	Sc PPM	Ta PPM	Ti PCT	Zr PPM	S PCT
26307		0.67	0.01	0.01	17	<1	<2	<1	<1	<5	<10	.014	1	<.01
26308		>10.00	<.01	0.02	1245	4	<2	10	<1	<5	<10	.038	2	0.33
26312		1.24	0.03	0.05	38	1	<2	7	<1	<5	<10	.021	<1	0.04
26313		>10.00	<.01	<.01	>2000	4	<2	3	<1	<5	<10	.022	2	0.15
26315		>10.00	<.01	0.01	>2000	5	<2	3	<1	<5	<10	.026	2	0.35
26316		7.04	0.03	0.04	90	17	<2	1	<1	<5	<10	.048	4	0.06
26317		>10.00	0.02	0.01	>2000	4	<2	2	<1	<5	<10	.016	<1	0.14